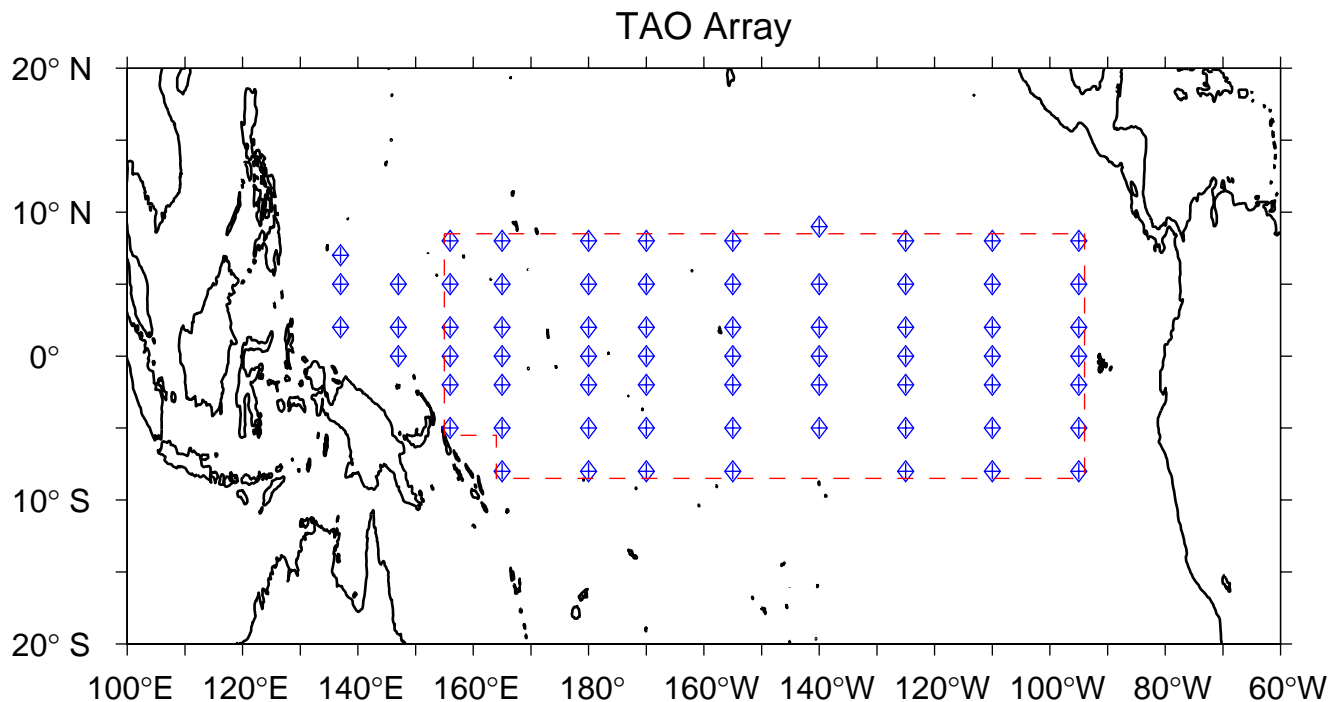


Warm water volume changes and transports in the tropical Pacific during the 1997–1998 El Niño

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Basic Premise

- The build up and decline of warm water volume in the equatorial Pacific is directly related to the timing of El Niño events.

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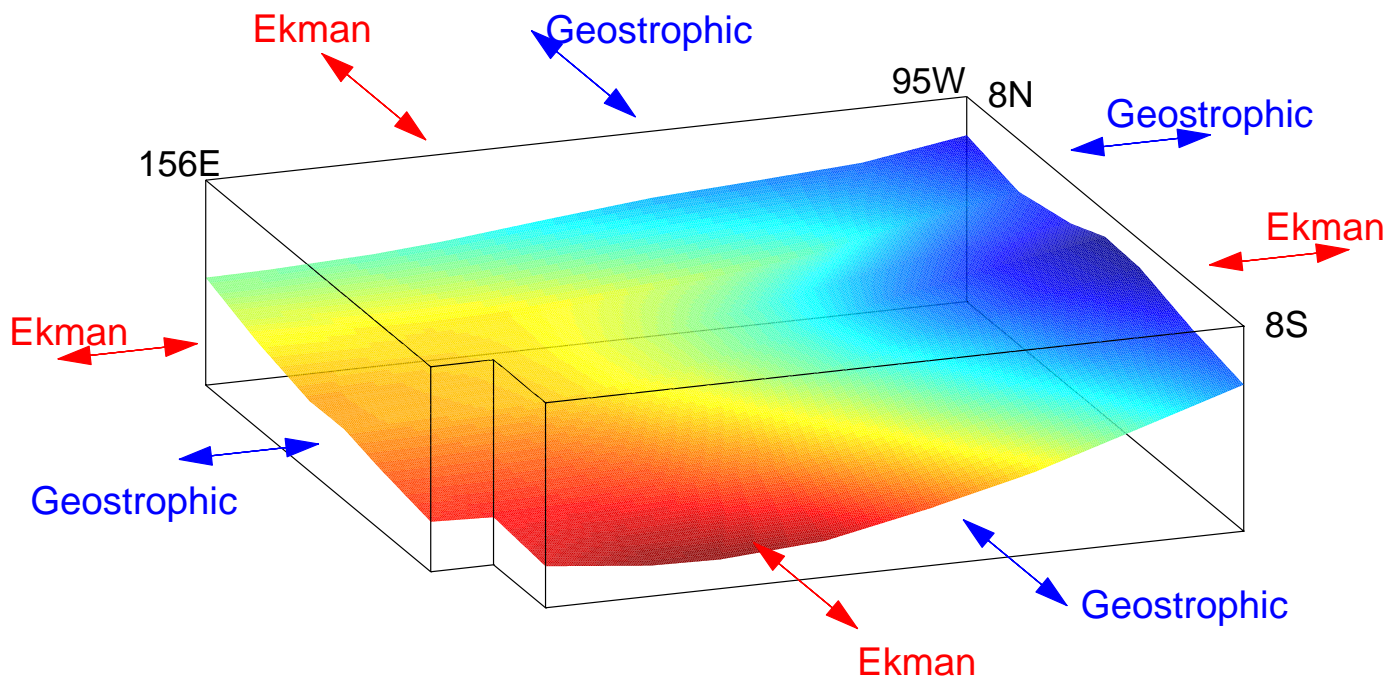
[§]NOAA/Pacific Marine Environmental Laboratory, Seattle, WA

Specific objectives

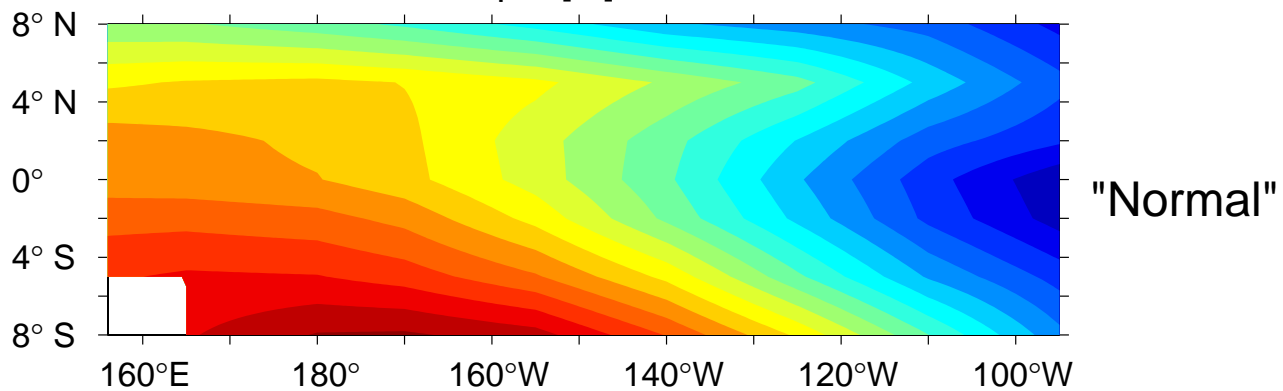
- To quantify the “flushing” of warm water to higher latitudes during the El Niño.
- To quantify the accumulation of warm water between the end of one El Niño and the beginning of the next.
- To quantify the lead time between the peak in the warm water volume and the peak in the SST.
- To compare the 1994 and 1997 El Niño signals and quantify the differences.

Advantages of this Study

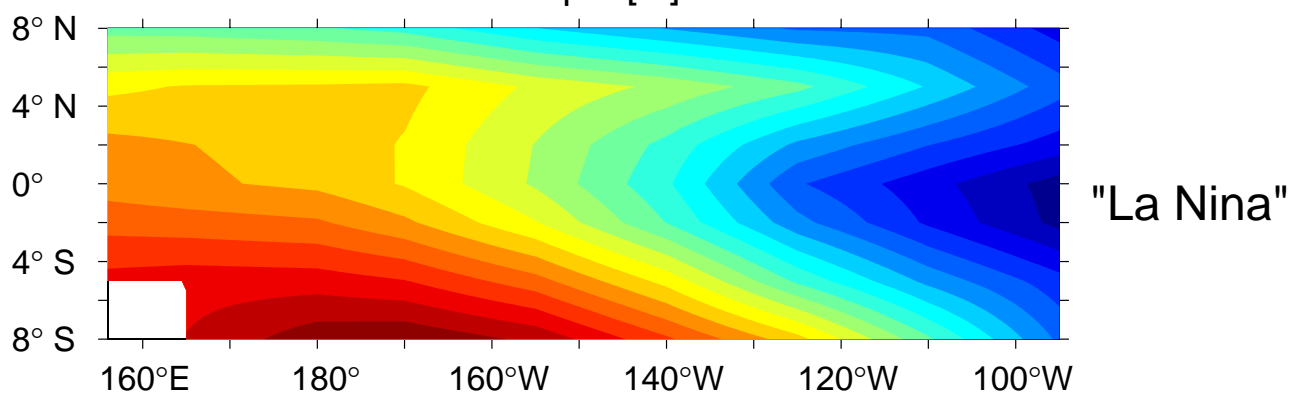
- More observations have been made during this El Niño than during any previous El Niño event.
- This El Niño event is one of the largest ever observed, resulting in a favorable signal-to-noise ratio.



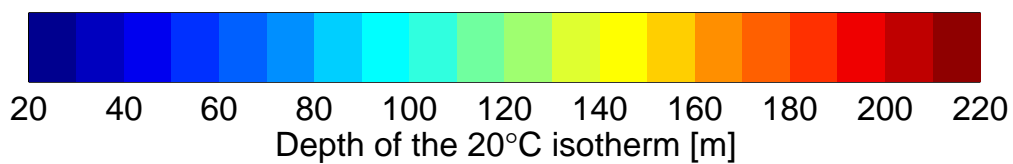
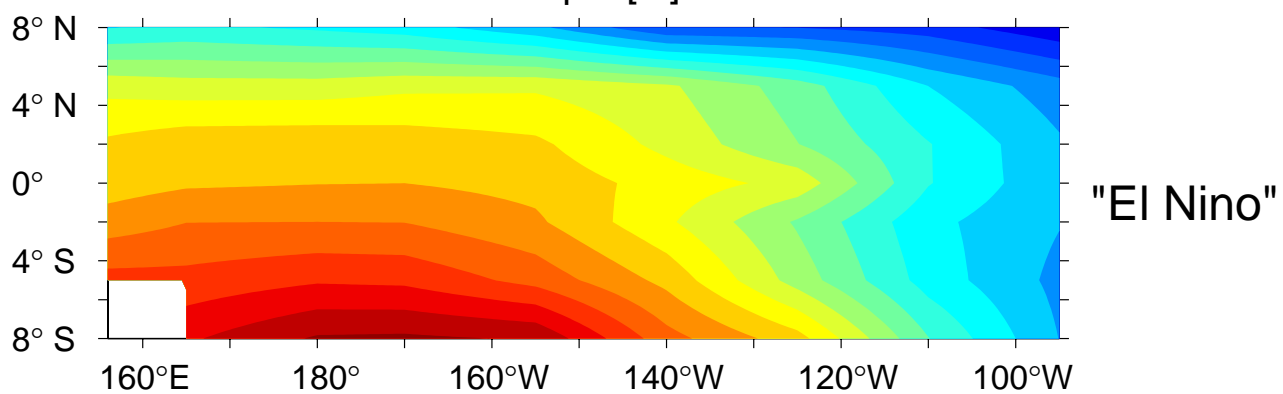
20°C isotherm depth [m]: 1993 & 1996 mean

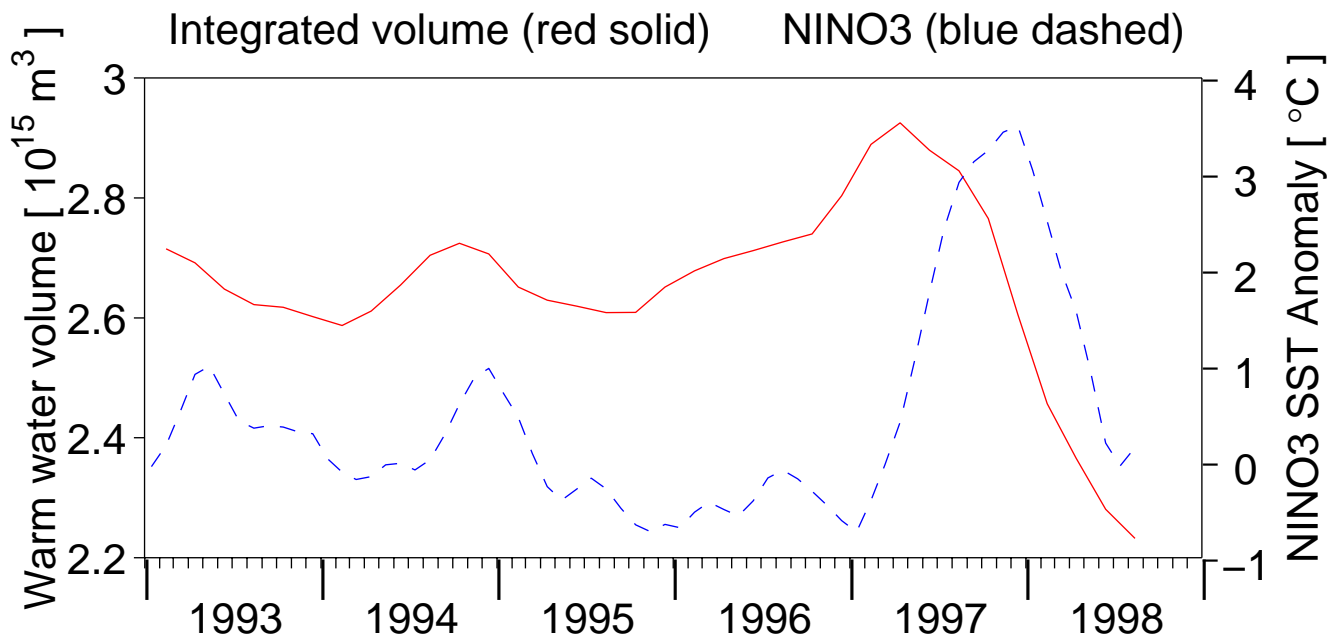


20°C isotherm depth [m]: 1995 mean

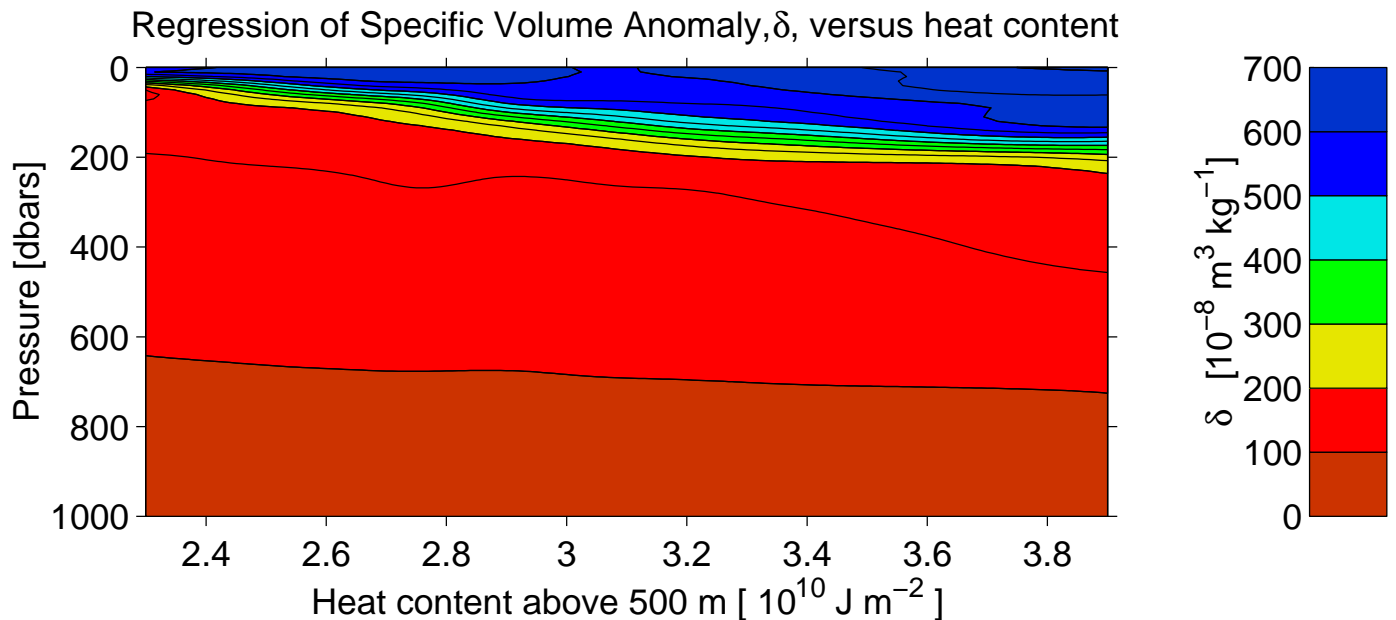


20°C isotherm depth [m]: 1997 mean



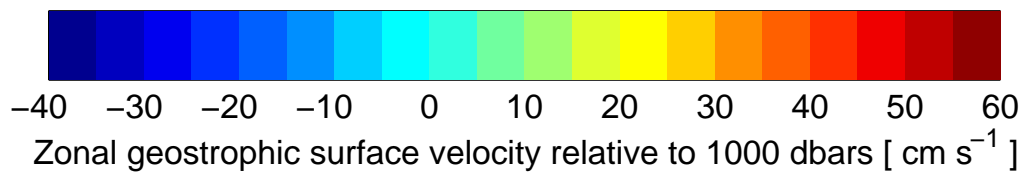
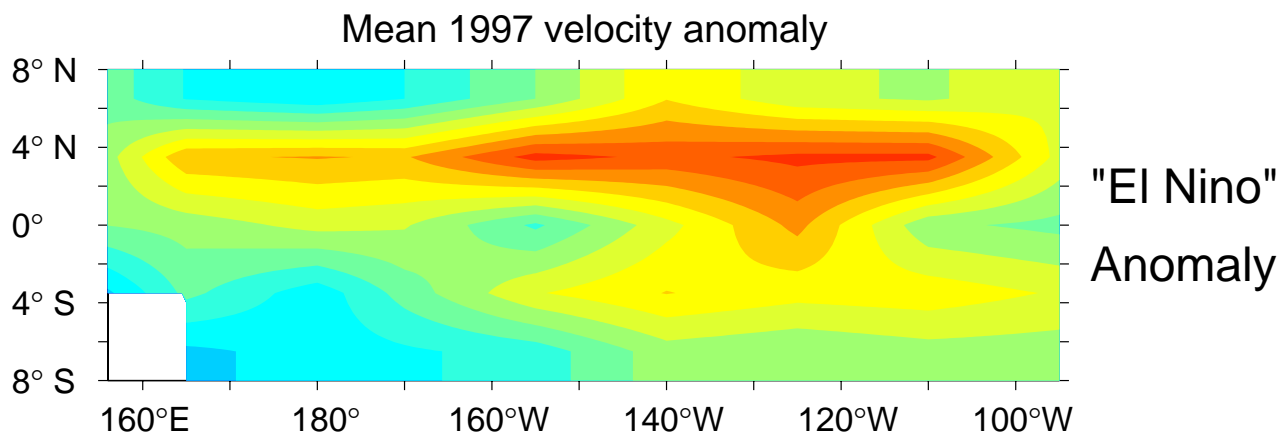
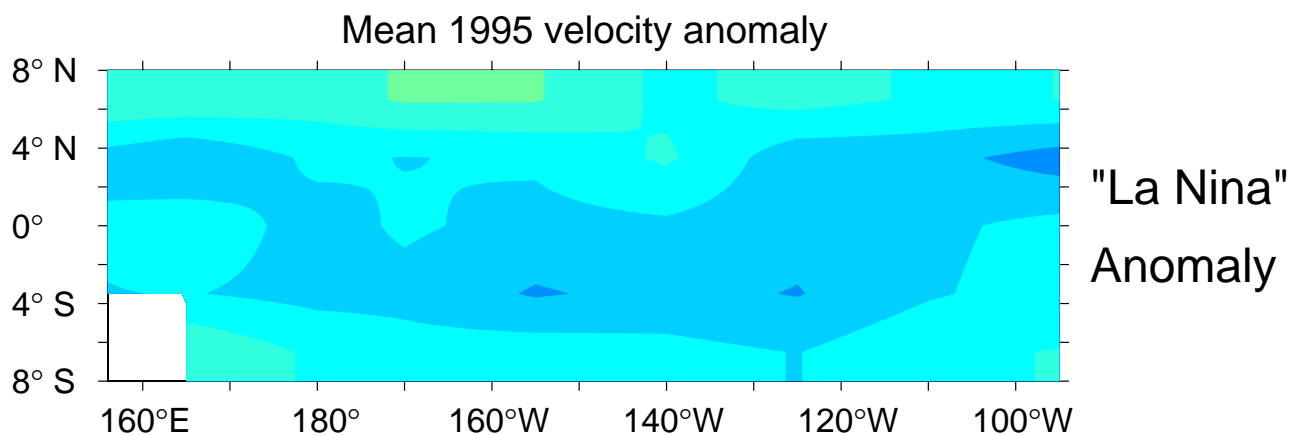
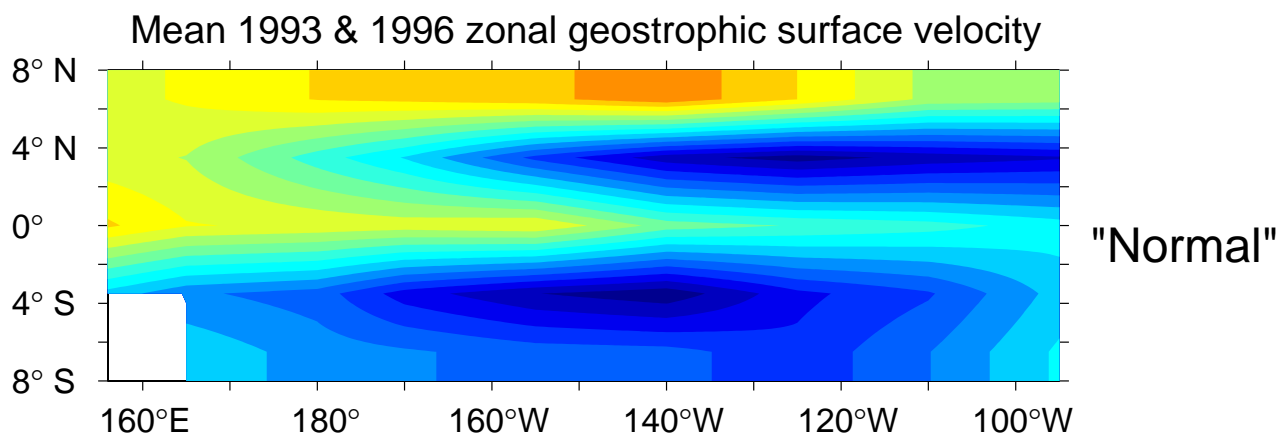


- Warm water volume decreases from Jan. 1993 until Feb. 1994.
- Volume increases from Feb. 1994 until Oct. 1994, this is the buildup prior to the 1994 El Niño.
- Volume decreases from Oct. 1994 until Oct. 1995, this period includes the 1994 El Niño and the beginning of the 1995 La Niña.
- Volume increases from Oct. 1995 until Apr. 1997, this is the buildup prior to the 1997 El Niño.
- Volume decreases from Apr. 1997 until Sep. 1998, this period includes the 1997 El Niño and the beginning of the 1998 La Niña.

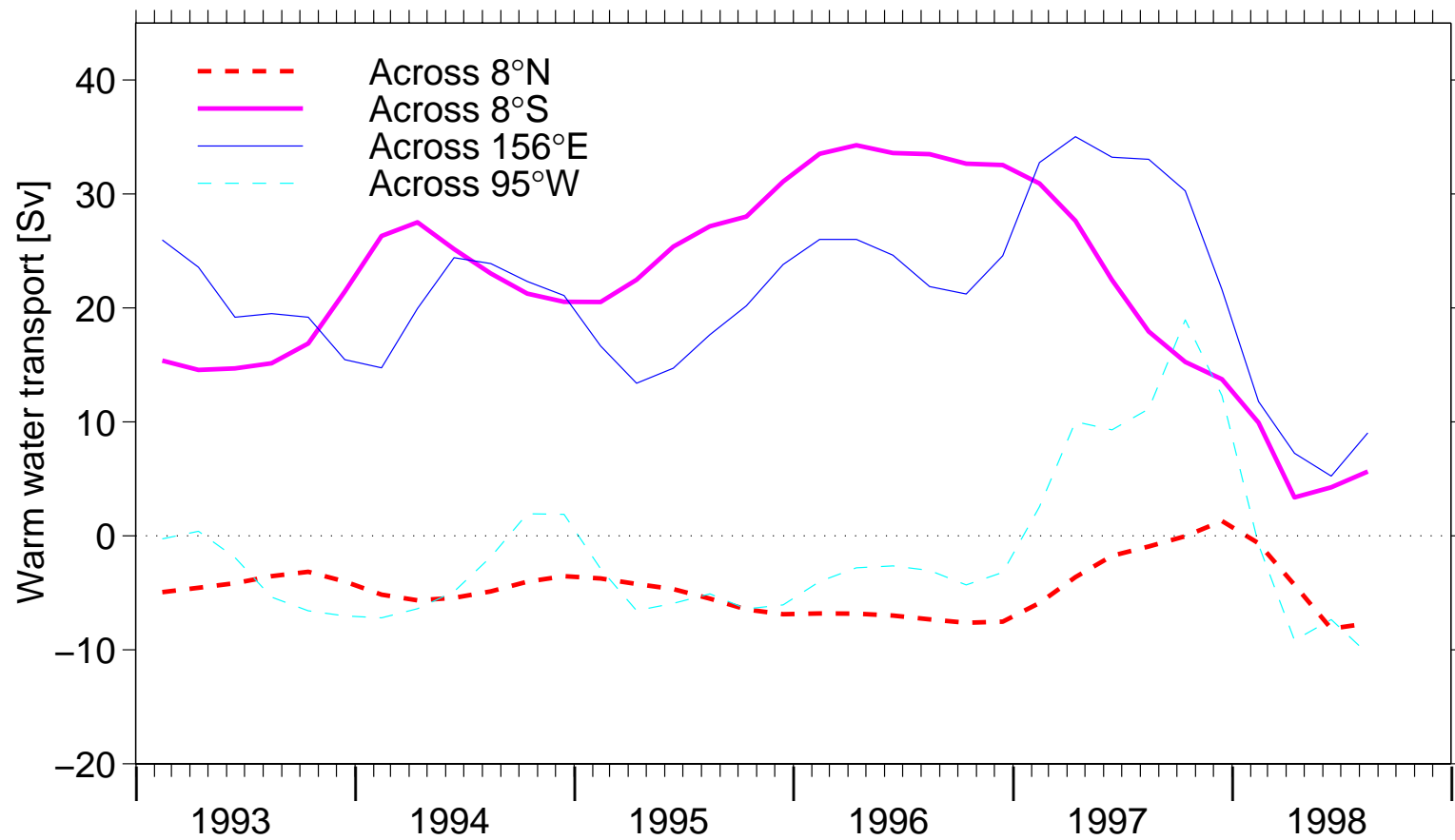


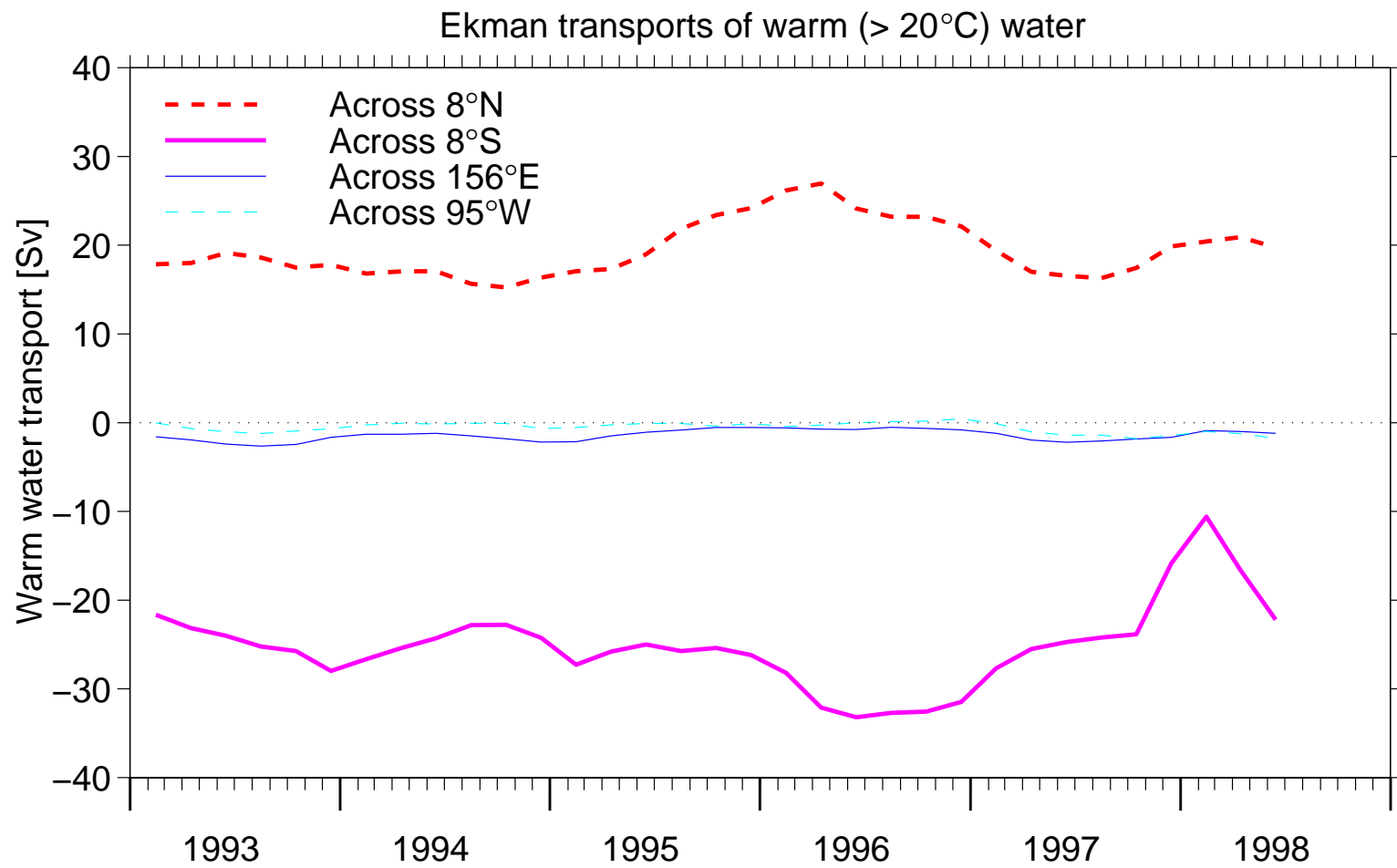
Determining velocity from TAO temperature measurements

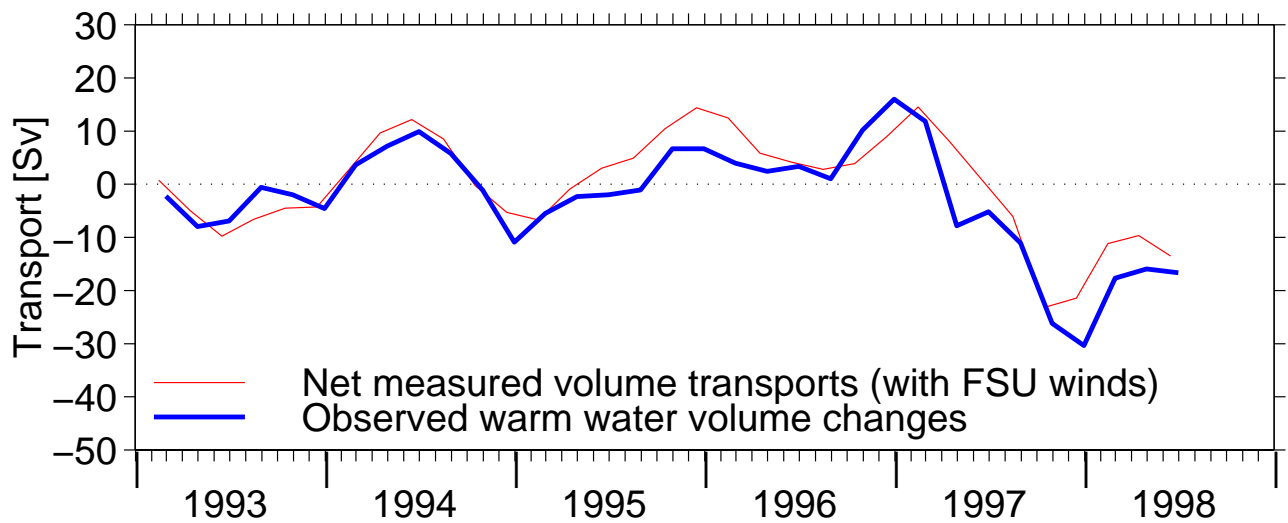
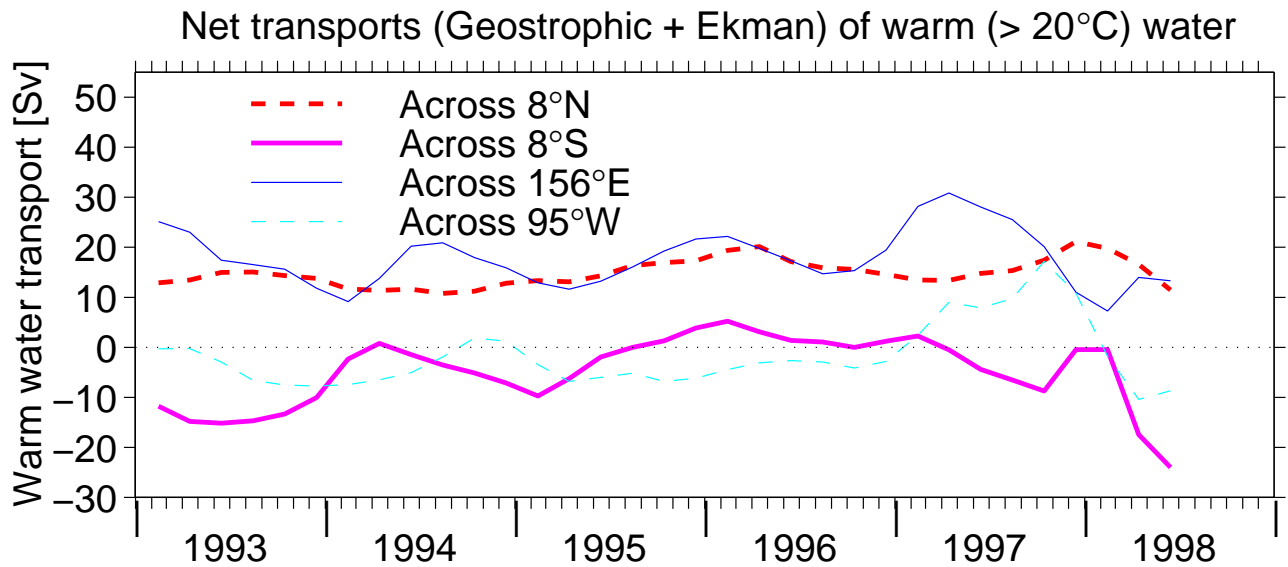
- Regress specific volume anomaly against heat content integrated above 500 dbar using historical hydrography (see above figure). [Meinen and Watts, submitted to JGR, 1998]
- Use timeseries of heat content from TAO buoys along with regression to give profile of specific volume anomaly at each mooring for each time step.
- Integrate profiles of specific volume anomaly to give geopotential height anomaly.
- Use the dynamic method to determine profiles of geostrophic velocity relative to 1000 dbars.



Geostrophic transports of warm ($> 20^{\circ}\text{C}$) water







- Correlation coefficient $r = 0.90$
Mean difference = -2.5 Sv
RMS difference = 5.1 Sv

Conclusions

- The volume of warm water showed a significant increase between the 1995 La Niña and the 1997 El Niño, consistent with the model results of Cane et al. and others.
- The peak in warm water volume preceded the peak in sea-surface temperature (integrated over the same region) by 6 months (Apr. 1997 vs. Oct. 1997).
- The loss of warm water during the 1997 El Niño was about 6 times larger than the loss during the 1994 El Niño.
- During the 1997 El Niño the region bounded by 156°E - 95°W , 8°S - 8°N lost about 28% of its warm water volume and about 31% of its heat content, confirming that one of the effects of El Niño is to transport considerable quantities of heat from the equator towards the poles.

Future work

- Extend the calculation to the eastern boundary and also westward up to the western boundary currents using TOPEX altimetry and XBT data.
- Complete a heat budget calculation for the same region.
- Determine the meridional extent of the El Niño warm water transports using TOPEX and other data sources.